## Background

Using linearized least-squares inversion (LLSI), the depth to the top of the low-velocity zone (LVZ) beneath oceans can be found along paths for which surface-wave dispersion is well constrained. The same techniques may give inconclusive results when applied to continental shield regions. The conclusion drawn by Snoke \& James (S\&J) in their LLSI for the S-wave structure beneath the eastern Paraná Basin in central Brazil is that there is no resolvable LVZ to a least 150 km depth. See Fig. 1 below.


Figure 1: Top: Four interstation paths in the Paranß Basin used for surface-wave analysis. (Three are between the same two stations.) Bottom left: fundamental-mode Rayleigh and Love Phase and Group velocities. Symbols are the data, and lines Each dispersion velocity datum is a composite from up to four events, weighted by the inverse of the estimated variances. Bottom right: Best-fit Swave model (LLSI-PAR) calculated using LLSI (solid) compared with the continental PEM model (dashed) for the mantle. The LLSI-PAR velocity model has 43 constant-velocity layers, and the damping used was 0.1 times the maximum
eigenvalue of the data kernel matrix.

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References: S&J: JGR, 102, 2939-2951, 1997
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## Neighbourhood Algorithm

Because of nonlinearities in the model-data relationships, the damping required to stabilize the inversion and smoothing constraints, the LLSI provides little easily interpreted information about model variance or resolution or the ensemble of acceptable models. Sambridge has introduced the Neighbourhood Algorithm (NA), a direct-search method for nonlinear inversion in a multidimensional parameter (model) space, which can be tuned to extract information from an acceptable ensemble of models in addition to finding a single best-fit model. Here we apply NA to the S\&J data set to see if it can provide more information about th velocity structure at depth beneath the eastern Parana Basin.


Figure 2: The three stages of a Neighbourhood Algorithm search. The top left panel shows an initial 10 uniformly randomly distributed points and their corresponding Voronoi cells. The top right points generated by the NA, and the bottom left panel shows the Voronoi cells for 500 points. The bottom right panel shows the true fitness landscape, darker shades are higher fitness. With increased sampling in NA, the concentration is much higher in the regions of higher fitness, and all four maxima are found by NA.

## Procedure

The model parameters used in this study are overlapping, weighted averages over the velocitydepth model (Fig. 3). The data misfit for each mode realization is the square of the length of the error vector with each element weighted by the inverse of that datums variance. An additional smoothing
constraint is imposed by adding 5.0 to a misfit if successive parameters from among parameters 2 through 8 have opposite signs and differ by more than a preset value ( 0.175 for these runs). For the two NA runs discussed here, the tuning parameters are $n_{\text {si }}=500, n_{s}=100, n_{r}=50$, and $N=95$ giving a total of 10,000 model evaluations. Herrmann's SURF routines are used for the forward modeling.


Figure 4: The full ensemble of 10,000 models produced by the NA for the Paranß dispersion data model is represented by a dot colour coded by data misfit, and the parameter ranges shown are $1 / 6$ the prescribed hard limits. The concentration of sampling increases in the regions of better data fit, but the full range of values between the hard limits are sample The correlations of data fit indicate that the crustal velocity (Vs Crust) is more tightly constrained tha ntle (Vsn).


Figure 5: Top: Predicted dispersions and models for the ensemble of all models from the NA run total of 10,000 evaluations which have a misfit less than 0.0111. Notation as in Fig. 1. Bottom: Average model and its dispersion calculated from the ensemble shown above. The misfit is 0.0107 , which would rank it as second among the full ensemble. Calculated (relative) standard deviations for the average velocity model are included

Figure 6: As shown in Fig. 5, the ensemble trend and the average model (labeled NA-PAR) diffe significantly from the reference model (labeled PEM-C42). Shown here are those two velocity models along with the model produced from a nn on a data set with dispersion velocities at the same periods and with the same errors as the data hown in Figs. 1 and 5, but the velocities were

S Velocity $\underset{4.0}{(\mathrm{~km} / \mathrm{s})}$
NA-PAR
PEM-C42
NA-PEM
NA-PAR
PEM-C42
NA-PEM
${ }^{4.8}$
Steps 2 and 3 constitute an iteration which is then repeated N times resulting in a total of $\mathrm{n}_{\mathrm{si}}+\mathrm{Nn}_{\mathrm{s}}$ model in Fig. 2.
Voronoi cells are nearest neighbour regions as defined by a distance norm. For any set of points (models) in a space with any namber of dimensions (unknowns), the Voronoi cells are unique, he model space into a series of neighborhoods.


fixed to give an exact fit for model PEM-C42. This shows that if there were a LVZ at depths shallower than 150 km , it could be resolved by our data set.

## Conclusions

Figure 3: Interpolation model parameters: Of the eight parameters, two are box car in shape resulting in uniform weighting of perturbations to the base-model (the LLSI-PAR model) throughout the depth range, and the remaining six are overlapping triangles. Ranges for these parameters increase from $\pm 0.6 \mathrm{~km} / \mathrm{s}$ for the crustal velocities (parameter 1) to $\pm 1.75 \mathrm{~km} / \mathrm{s}$ for parameter 8 .

Applying a Neighbourhood Algorithm analysis to dispersion velocity data provides insights not easily seen from LLSI analyses. It is easy to test the relative importance from among the data by doing NA runs with data subsets. By varying the parameterization scheme or the definition of misfit, one can concentrate the analysis on different parts of the model space. For this application, NA confirms the LLSI conclusion that no LVZ begins shallower than 150 km depth but adds the insight that there is a negative gradient in velocity below 150 km depth. To constrain the velocity structure at greater depths requires dispersion data at higher periods.

